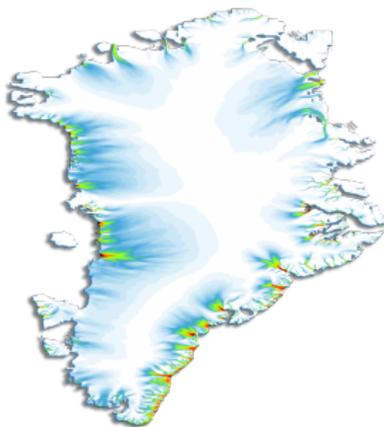


# Using observations to validate ice sheet models

Andy Aschwanden



# Outline

## Setting the stage

Thermodynamics

Boundary conditions

Model validation

# Sea Level Response to Ice Sheet Evolution

- ▶ “SeaRISE” led by Bob Bindschadler
- ▶ national and international **unfunded** participants

Journal of Glaciology, Vol. 59, No. 214, 2013 doi:10.3189/2013jog12J125

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## Ice-sheet model sensitivities to environmental forcing and their use in projecting future sea level (the SeaRISE project)

Robert A. BINDSCHADLER,<sup>1</sup> Sophie NOWICKI,<sup>1</sup> Ayako Abe-OUCHI,<sup>2</sup>  
Andy ASCHWANDEN,<sup>3</sup> Hyeungu CHOI,<sup>4</sup> Jim FASTOOK,<sup>5</sup> Glen GRANZOW,<sup>6</sup>  
Ralf GREVE,<sup>7</sup> Gail GUTOWSKI,<sup>8</sup> Ute HERZFELD,<sup>9</sup> Charles JACKSON,<sup>8</sup>  
Jesse JOHNSON,<sup>6</sup> Constantine KHROULEV,<sup>3</sup> Anders LEVERMANN,<sup>10</sup>  
William H. LIPSCOMB,<sup>11</sup> Maria A. MARTIN,<sup>12</sup> Mathieu MORLIGHEM,<sup>13</sup>  
Byron R. PARIZEK,<sup>14</sup> David POLLARD,<sup>15</sup> Stephen F. PRICE,<sup>11</sup> Diandong REN,<sup>16</sup>  
Fuyuki SAITO,<sup>17</sup> Tatsuru SATO,<sup>7</sup> Hakime SEDDIK,<sup>7</sup> Helene SEROUSSI,<sup>18</sup>  
Kunio TAKAHASHI,<sup>17</sup> Ryan WALKER,<sup>19</sup> Wei Li WANG<sup>1</sup>

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<sup>12</sup>Potsdam Institute for Climate Impact Research, Potsdam, Germany

<sup>13</sup>Department of Earth System Science, University of California, Irvine, Irvine, CA, USA

<sup>14</sup>Mathematics and Geoscience, Penn State DuBois, DuBois, PA, USA

<sup>15</sup>Earth and Environmental Systems Institute, The Pennsylvania State University, University Park, PA, USA

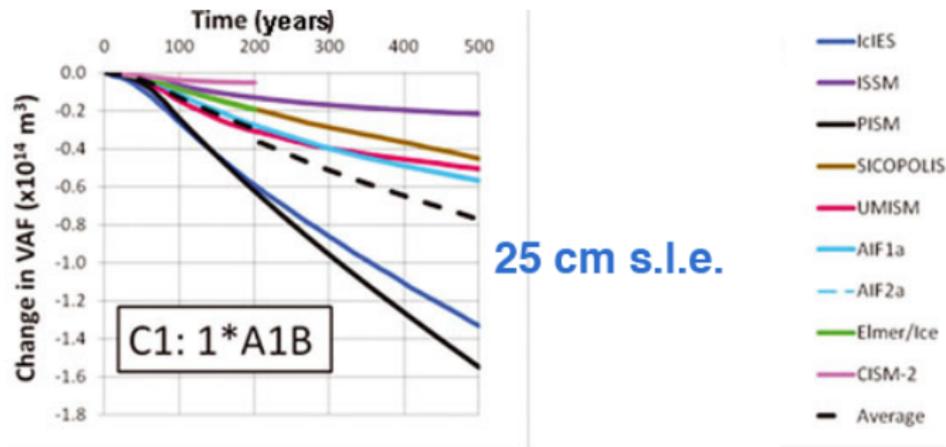
<sup>16</sup>Department of Physics, Curtin University of Technology, Perth, Australia

<sup>17</sup>Japan Agency for Marine–Earth Science and Technology, Research Institute for Global Change, Showamachi, Kanazawa, Yokohama, Kanagawa, Japan

<sup>18</sup>Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, USA

<sup>19</sup>Earth System Science Interdisciplinary Center, University of Maryland, College Park, MD, USA

# “SeaRISE”



Bindschadler et al. (2013), mod.

Would you rather trust

- ▶ a particular model
- ▶ the ensemble average
- ▶ none

⇒ let's look behind the scenes

# Parallel Ice Sheet Model

Documentation: [www.pism-docs.org](http://www.pism-docs.org)

Source code: <https://github.com/pism/pism>



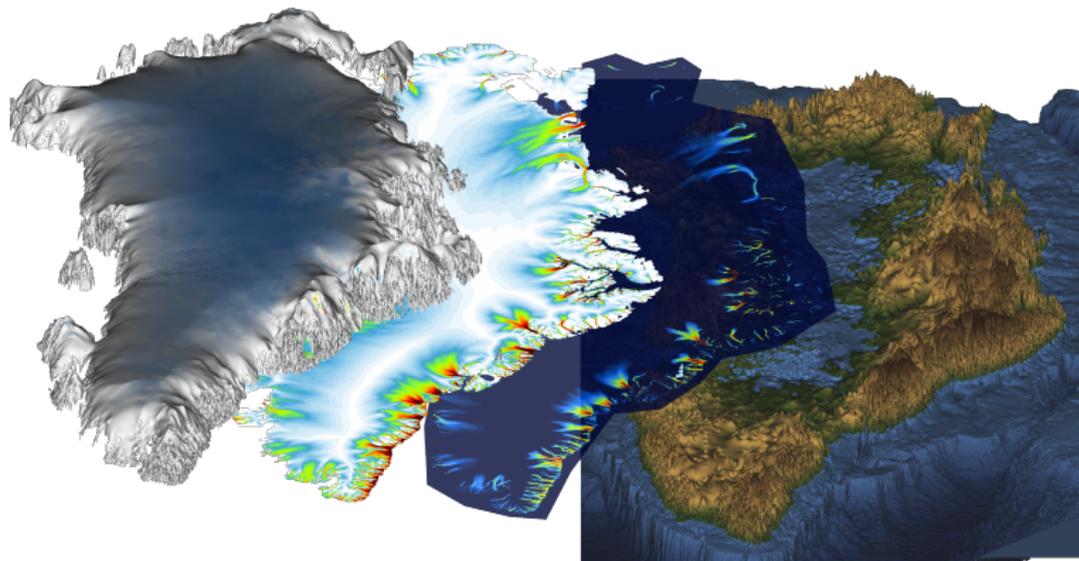
- ▶ open-source
- ▶ parallel
- ▶ high-resolution

- ▶ led by PI Ed Bueler, UAF
- ▶ jointly developed by UAF and Potsdam Institute for Climate Impact Research
- ▶ main software engineer: Constantine Khroulev, UAF
- ▶ > 20 contributors and users worldwide
- ▶ funded by



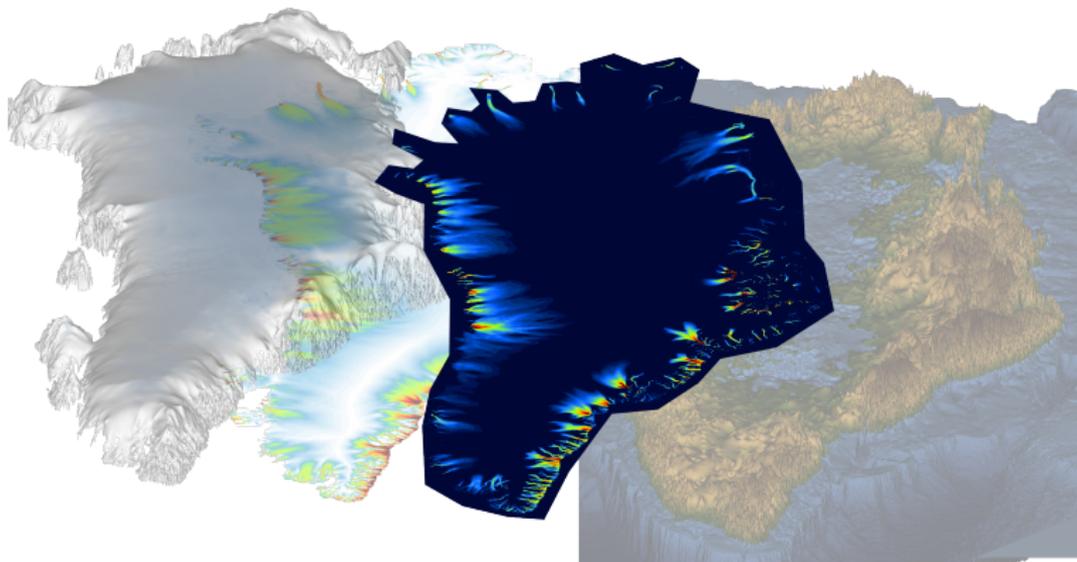


# Ice sheet system



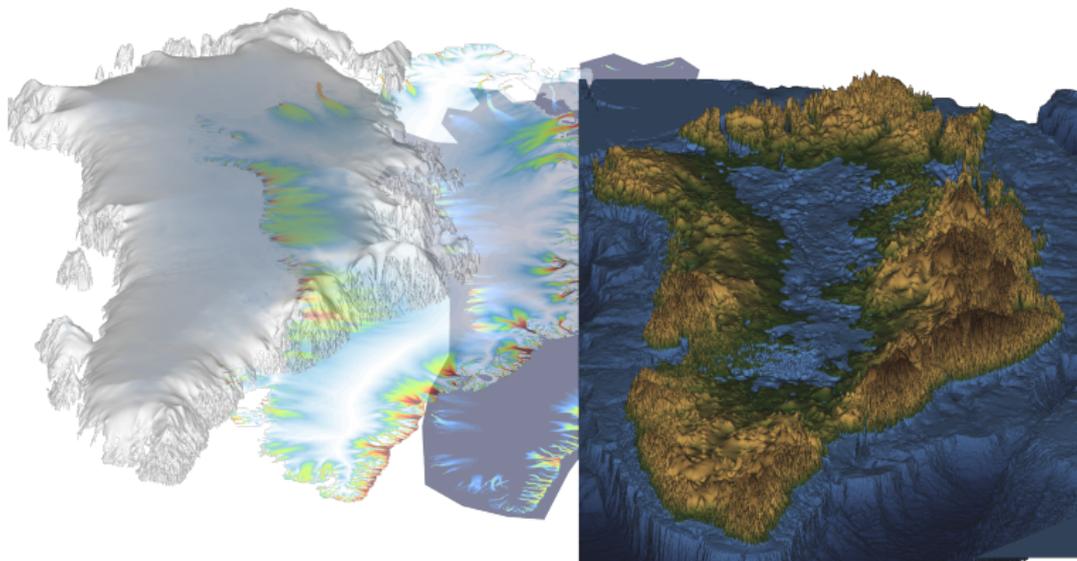
- ▶ ice dynamics
- ▶ thermodynamics
- ▶ surface processes
- ▶ boundary conditions
- ▶ hydrology
- ▶ ice-ocean interaction (e.g. calving)

# Ice sheet system



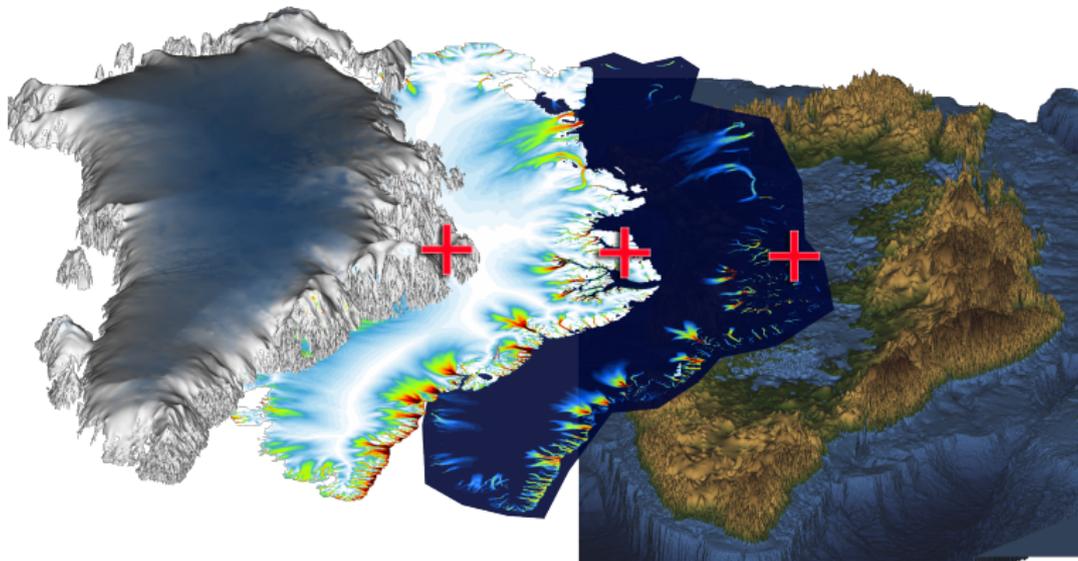
- ▶ ice dynamics
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# Ice sheet system



- ▶ ice dynamics
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# Ice sheet system



- ▶ ice dynamics
- ▶ thermodynamics
- ▶ surface processes
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- ▶ hydrology
- ▶ ice-ocean interaction (e.g. calving)

# Outline

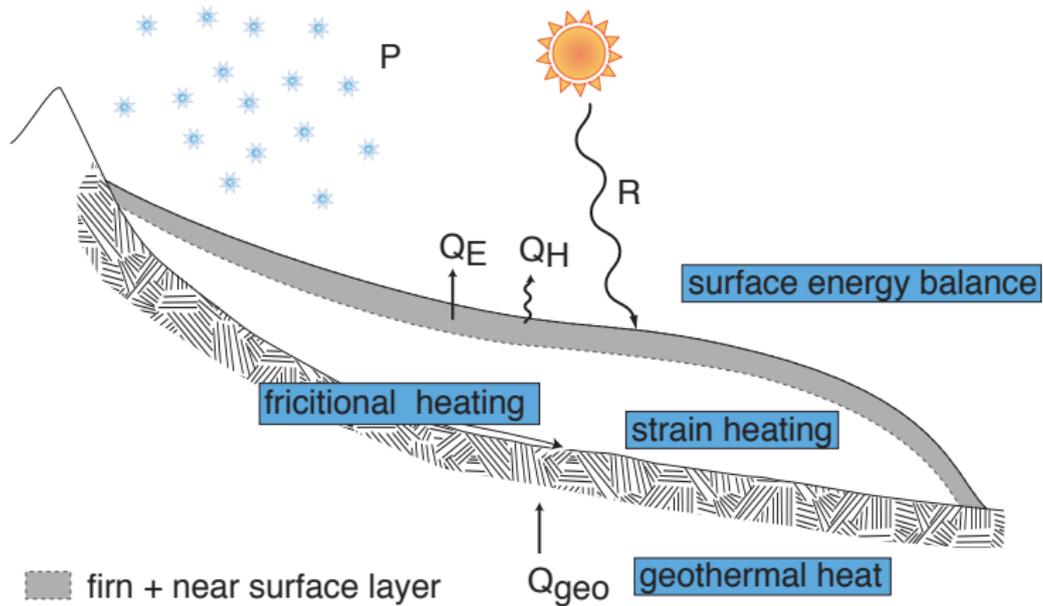
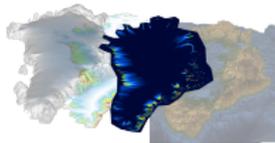
Setting the stage

**Thermodynamics**

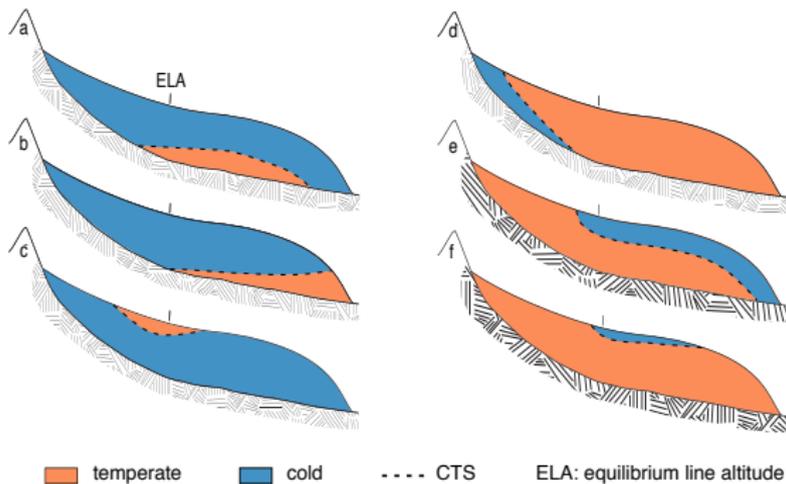
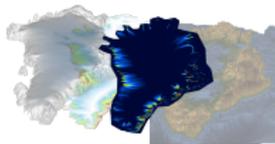
Boundary conditions

Model validation

# Heat sources



# Polythermal glaciers



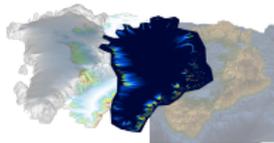
## Cold ice

- ▶ below pressure melting point

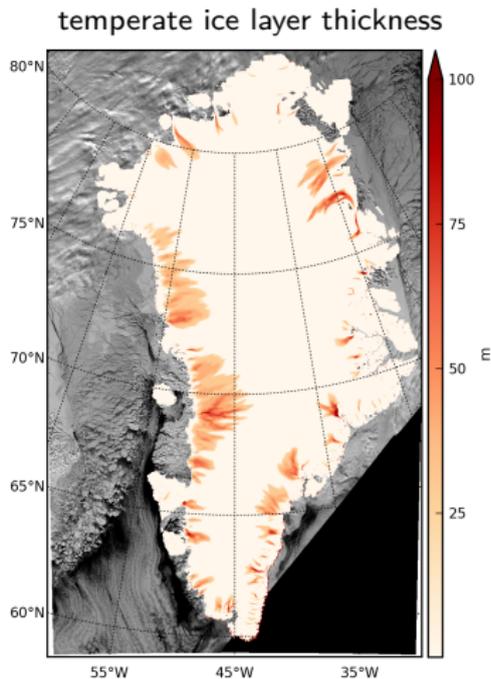
## Temperate ice

- ▶ at pressure melting point

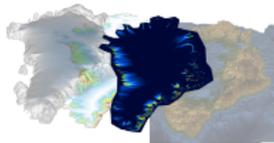
# Temperate ice in Greenland



- ▶  $\approx 35\%$  of the base is temperate (by area)
- ▶  $\approx 0.5\%$  of ice is temperate (by volume)
- ▶ but temperate where strain rates are already high



Aschwanden et al. (2012, modified)

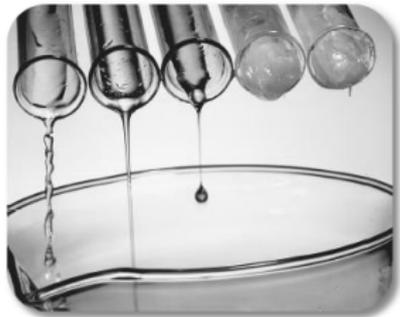


## Cold ice

- ▶ below pressure melting point
- ▶ solid phase only
- ▶ no liquid water content

## Temperate ice

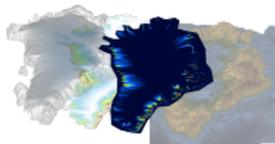
- ▶ at pressure melting point
- ▶ binary mixture of solid and liquid phase
- ▶ up to 5% liquid water within the ice matrix



Viscosity of ice depends on

- ▶ temperature
- ▶ liquid water fraction
- ▶ effective strain rate
- ▶ crystal orientation, impurities, etc.

# Enthalpy equation



- ▶ Conventional firn and glacier models are not energy conserving
- ▶ We replace the advection-diffusion-production equation for temperature with a similar equation for enthalpy (i.e. inner energy)

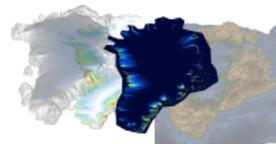
$$\rho \frac{\partial T}{\partial t} + \mathbf{v} \cdot \nabla T = -\nabla \cdot \mathbf{q} + Q$$

The equation above shows the standard heat balance equation. In the image, the terms  $\partial T$  and  $\nabla T$  are crossed out with red diagonal lines. A red  $H$  is placed above the first  $\partial$  and another red  $H$  is placed above the  $\nabla$ , indicating that the equation should be modified to use enthalpy  $H$  instead of temperature  $T$ .

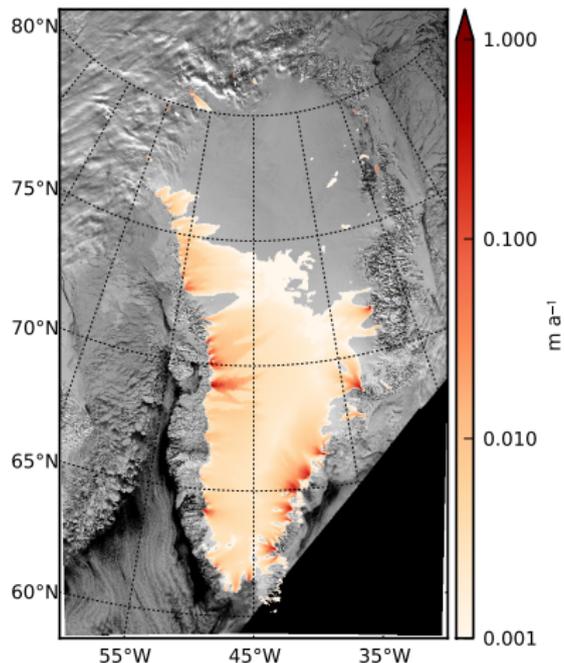
- ▶ Same PDE  $\Rightarrow$  relatively easy to implement

Aschwanden and Blatter (2009), Aschwanden et al. (2012)

# Basal melt rates



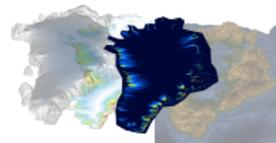
temperature equation



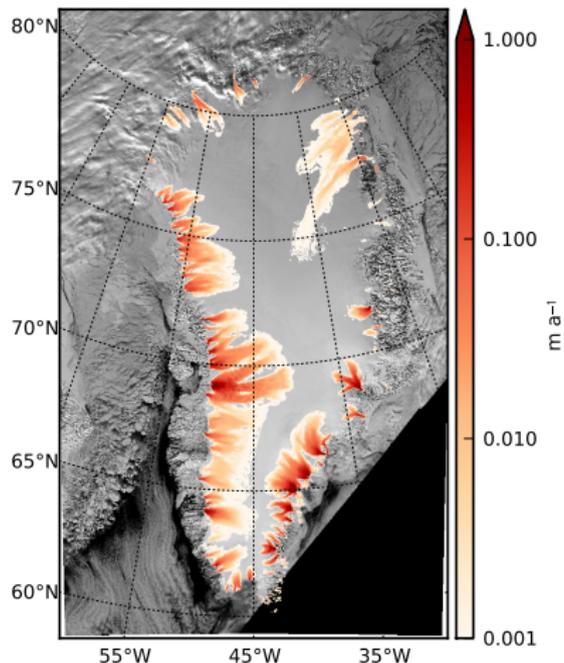
Aschwanden et al. (2012, modified)

- ▶ conservation of energy
- ▶ more realistic basal melt rates
- ▶ more realistic ice streams

# Basal melt rates



enthalpy equation



Aschwanden et al. (2012, modified)

- ▶ conservation of energy
- ▶ more realistic basal melt rates
- ▶ more realistic ice streams

# Outline

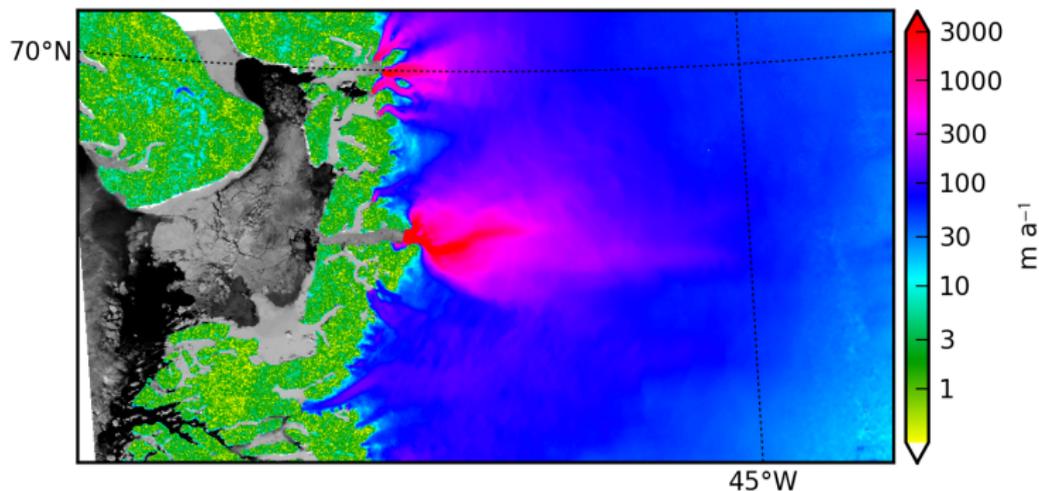
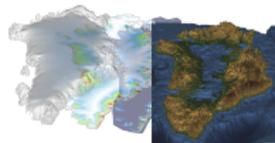
Setting the stage

Thermodynamics

**Boundary conditions**

Model validation

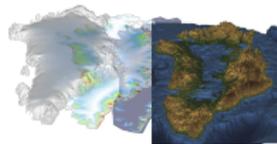
# Jakobshavn flows fast



Joughin et al. (2010)

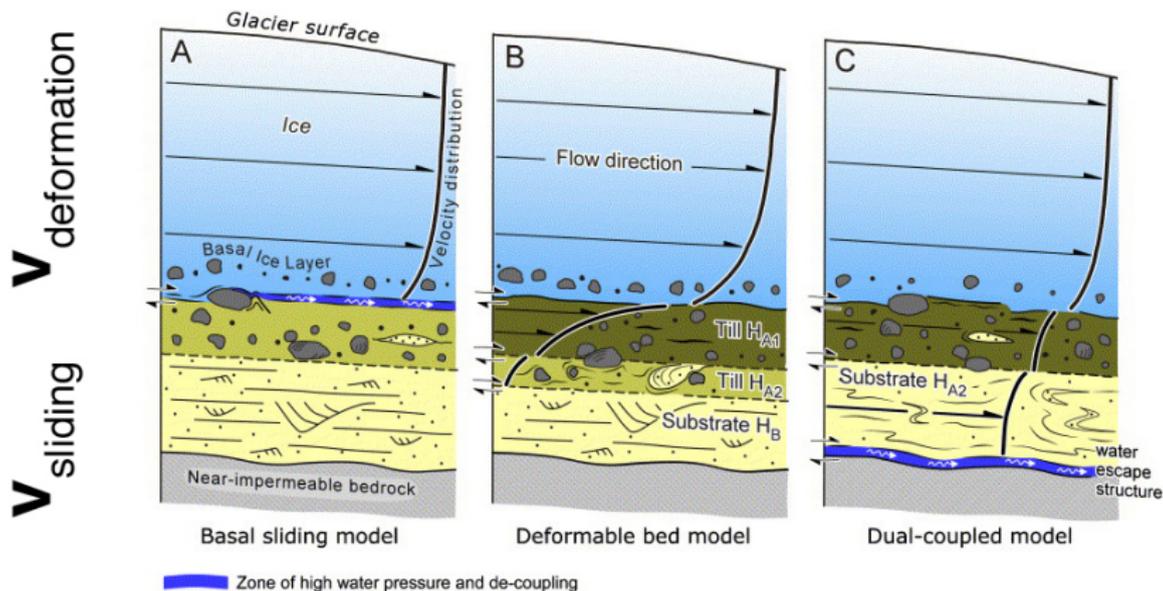
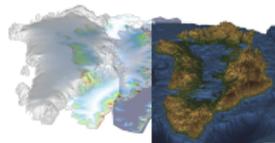
- ▶ Why does Jakobshavn flow so fast?
- ▶ not super exciting from above

# Jakobshavn flows fast



- ▶ Why does Jakobshavn flow so fast?
- ▶ not super exciting from above

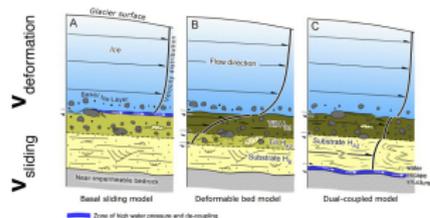
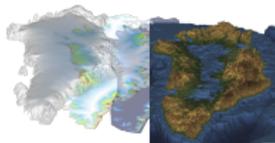
# Ice flow



Kjær et al., (2006, modified)

$$\mathbf{v} = \mathbf{v}_{\text{deformation}} + \mathbf{v}_{\text{sliding}}, \quad \mathbf{v} : \text{velocity}$$

# Ice flow



scaling arguments tell us:

$$\mathbf{v}_{\text{deformation}} \sim (\sin \alpha)^3 (H)^4$$

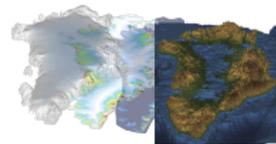
$H$ : ice thickness

$\alpha$ : surface slope

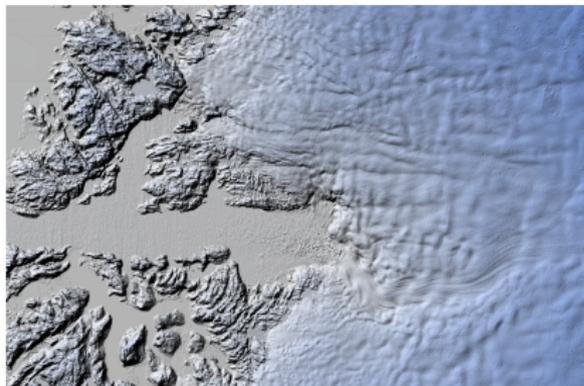
## Example

$$\frac{\delta H = 100 \text{ m}}{H = 1000 \text{ m}} = 10 \% \quad \Rightarrow \quad \frac{\delta \mathbf{v}}{\mathbf{v}} = 40 \%$$

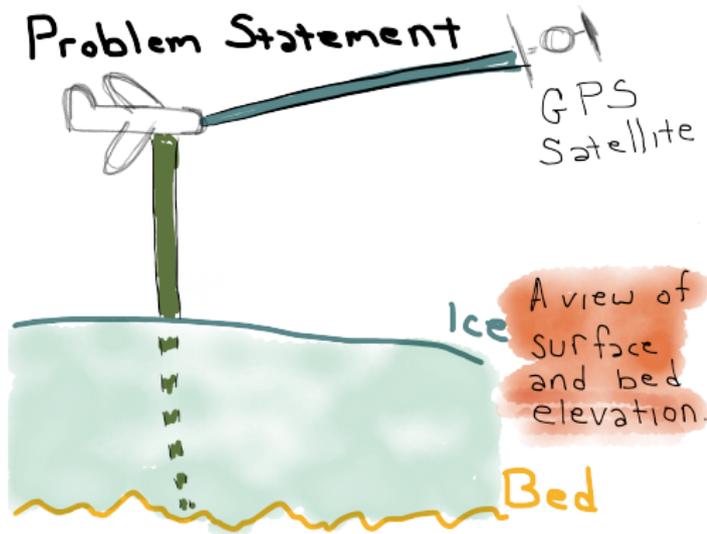
# Surface slope & ice thickness



surface elevation (slope)

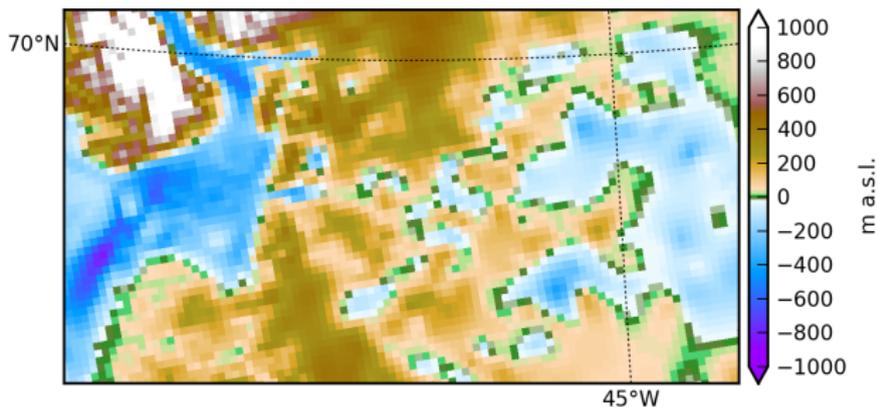
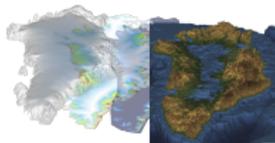


GIMP DEM (Howat et al.)



credit: Jesse Johnson

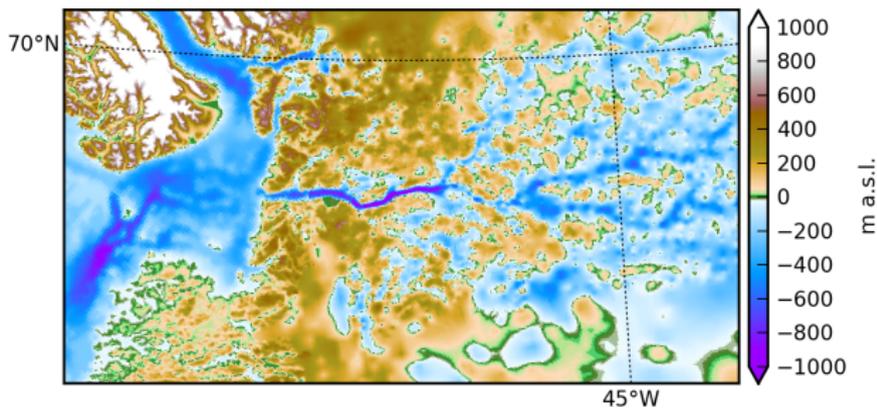
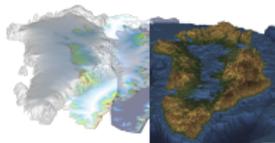
# Basal topography



Bamber et al. (2001)

- ▶ Operation Ice Bridge Mission since 2009
- ▶ Center for Remote Sensing (CReSIS) radar
- ▶ huge progress between 2001 and 2012

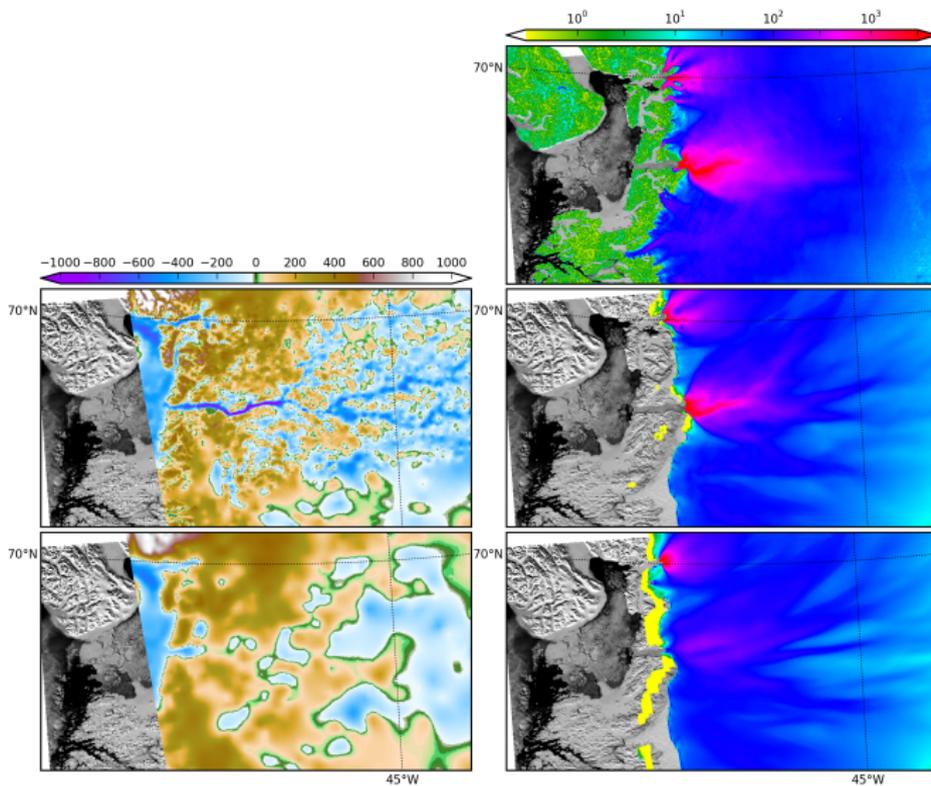
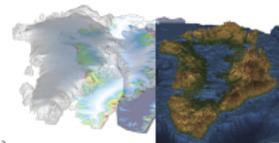
# Basal topography



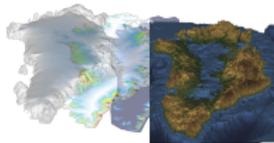
Griggs et al. (2012)

- ▶ Operation Ice Bridge Mission since 2009
- ▶ Center for Remote Sensing (CRSIS) radar
- ▶ huge progress between 2001 and 2012

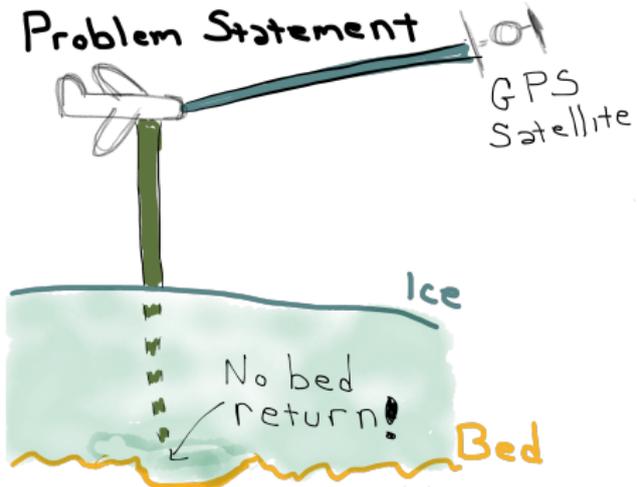
# It makes a difference



# Ice thickness

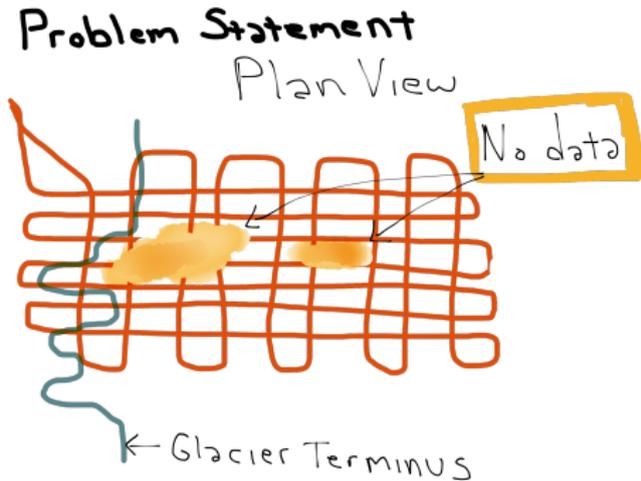


Problem Statement



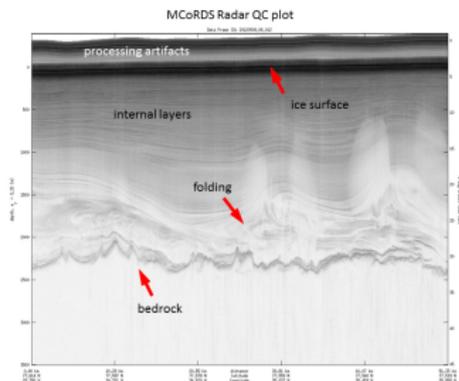
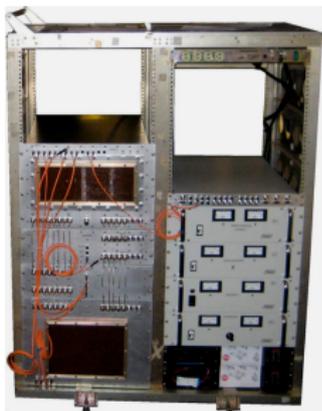
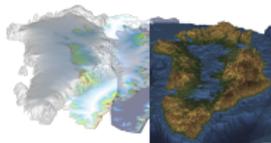
credit: Jesse Johnson

Problem Statement



credit: Jesse Johnson

# A new depth sounder



## MRI

Development of a high power, large-antenna array for a Basler for sounding and imaging of fast-flowing glaciers and ultra wideband radars to map near-surface internal layers. PI: Rick Hale, University of Kansas. NSF. Current support 2012–2014.

# Outline

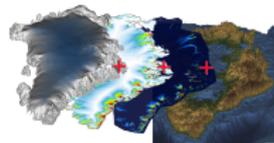
Setting the stage

Thermodynamics

Boundary conditions

**Model validation**

# Ice sheet model validation



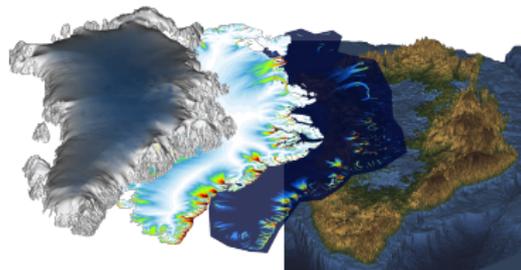
- ▶ comparing model results to a set of observations adequate to falsify a model

## Direct validation

of substantial sub-systems such as

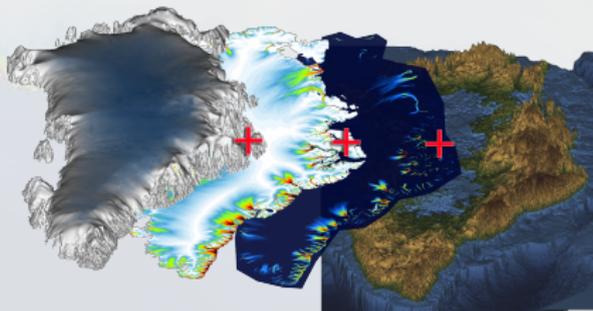
- ▶ basal hydrology
- ▶ thermodynamics
- ▶ ice dynamics

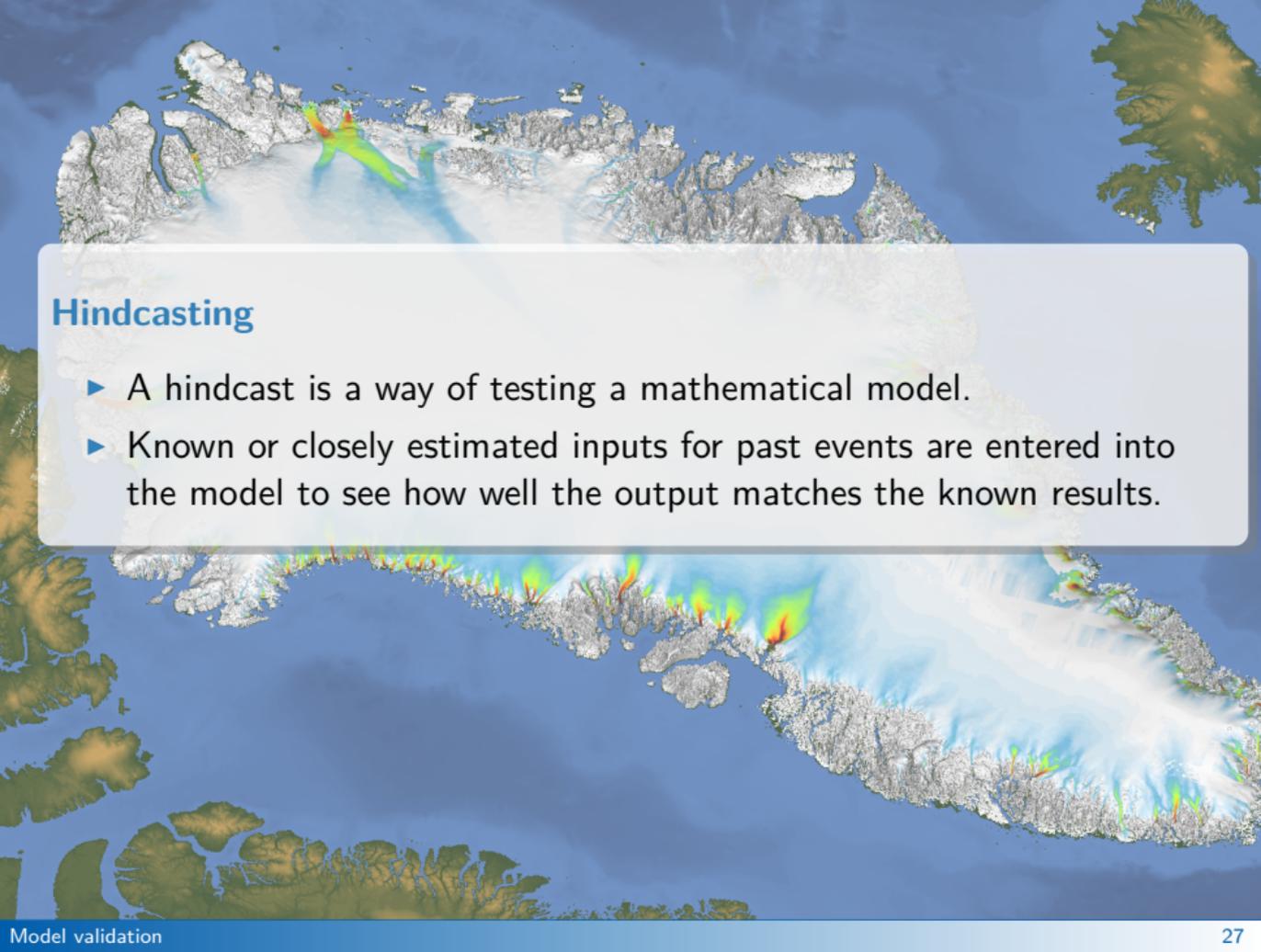
is **difficult** or **impossible**



## View as part of an earth system model

- ▶ we evaluate how the *system* responds to a given forcing
- ▶ “How successful is a state-of-the art ice sheet system model (i.e. the combination of physical models, their numerical approximations and implementations, and particular choices of boundary forcing and initial states) in reproducing observations of quantities such as ice thickness, and their temporal changes?”

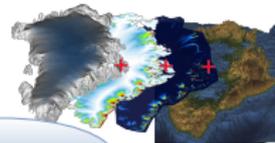




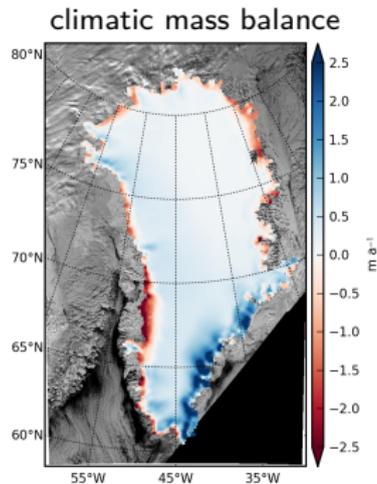
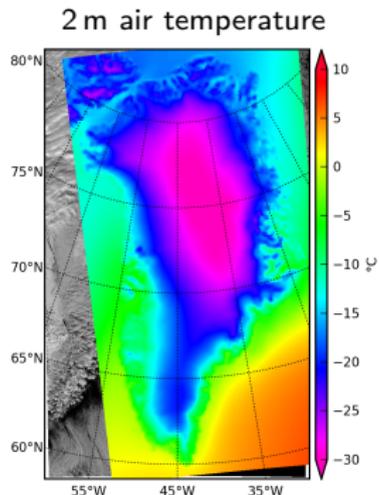
## Hindcasting

- ▶ A hindcast is a way of testing a mathematical model.
- ▶ Known or closely estimated inputs for past events are entered into the model to see how well the output matches the known results.

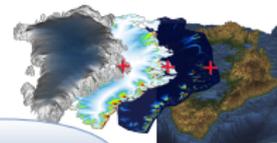
# Initialization



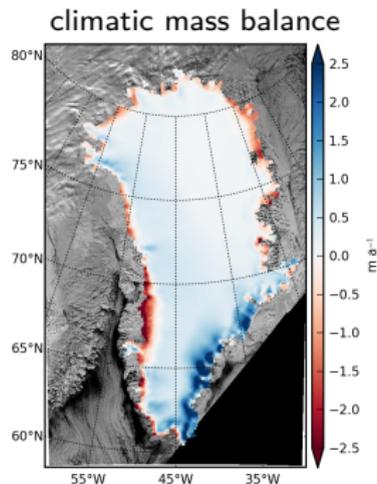
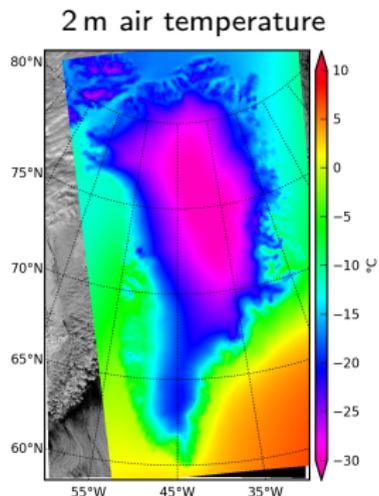
- ▶ RACMO2/GR driven by
  - ▶ ERA-reanalysis from 1961-2004
  - ▶ HadGEM2 from 1971-2004
- ▶ PISM driven by **mean values** of:



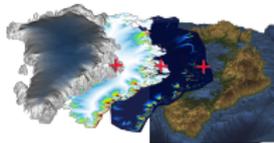
# Hindcast



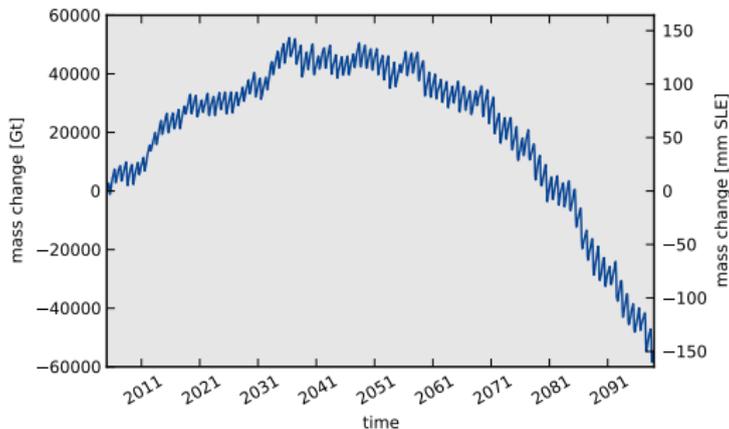
- ▶ RACMO2/GR driven by
  - ▶ ERA-reanalysis from 1961-2004
  - ▶ HadGEM2 from 1971-2004
- ▶ PISM driven by **monthly time-series** of:



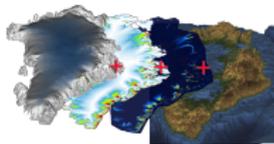
# Forecast



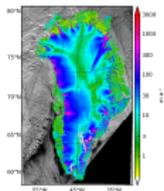
- ▶ Not the topic of this talk



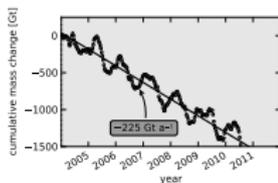
# Comparison with observations



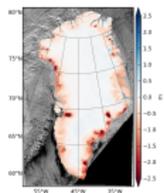
Hindcasts cover an era where we have a variety of in-situ and remotely-sensed observations such as:



- ▶ mean flow speed from 2000, 2006–2008 (SAR) from *Joughin et al.* (2010)

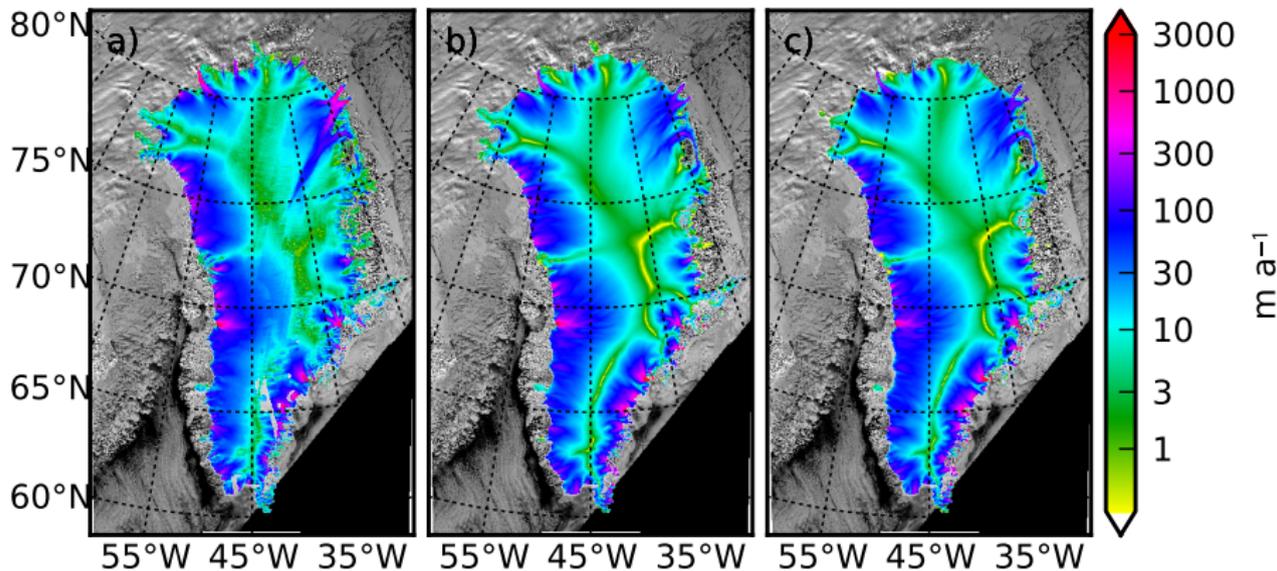
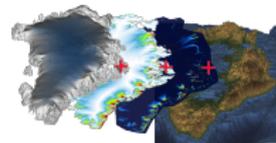


- ▶ cumulative mass change from 2003–2011 (GRACE) from *Luthcke et al.* (under review)



- ▶ elevation change from 2003–2009 (ICESat) from *Sørensen et al.* (2011)

# Flow speed



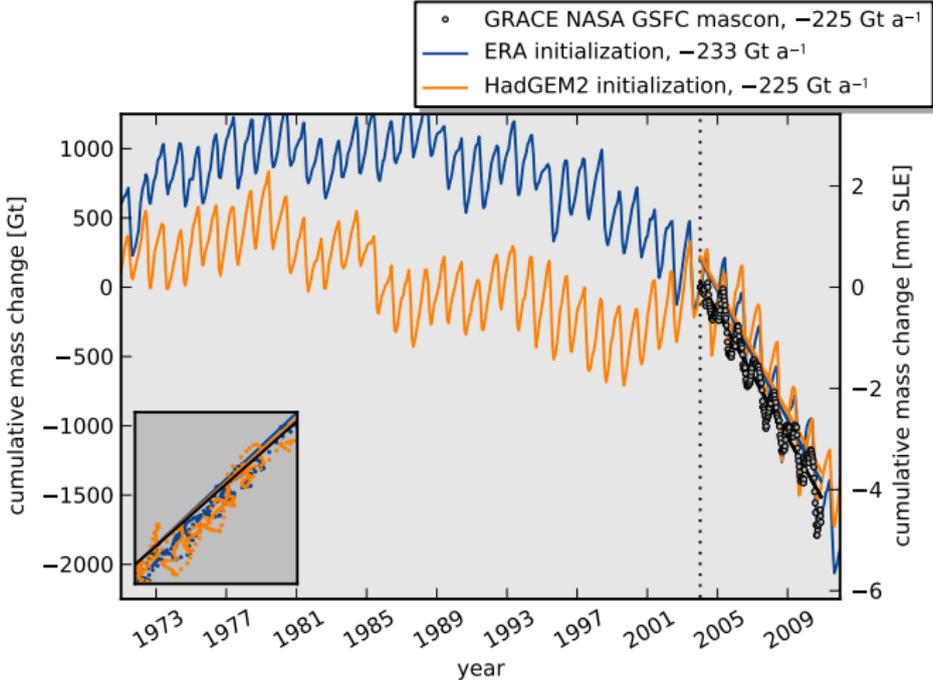
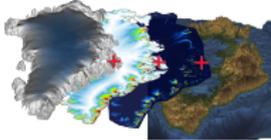
(a) SAR  
(Joughin et al, 2010)

(b) ERA init.

(c) HadGEM2 init.

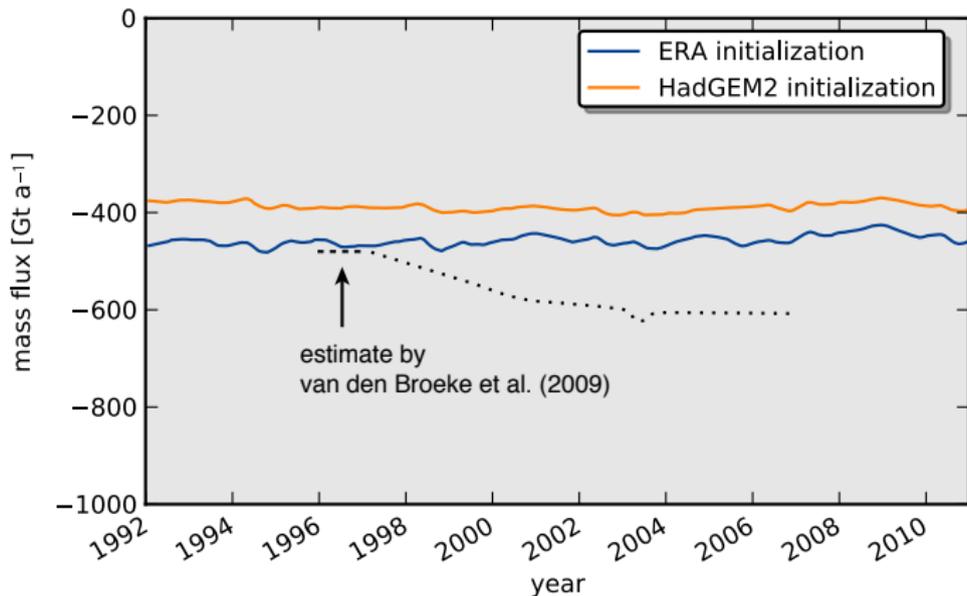
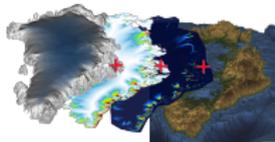
- ▶ reasonable agreement with observations

# Mass changes



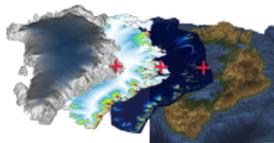
► an almost perfect fit (?)

# Ice discharge at ice/ocean interface

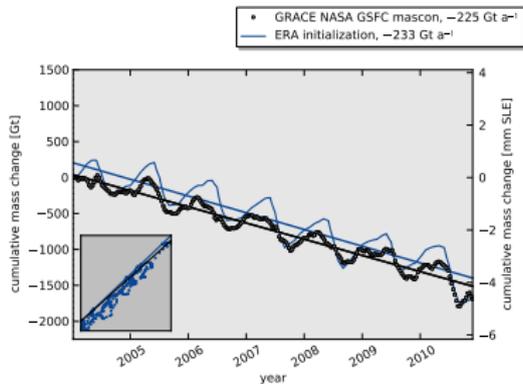


- ▶ simulated ice discharge remains nearly constant
- ▶ observed increase not simulated

# Wait a minute...

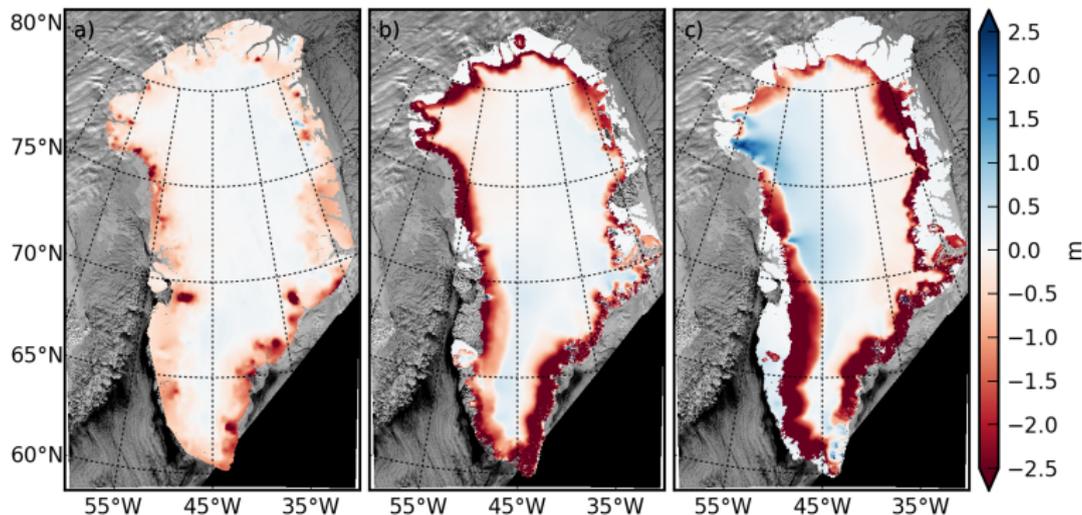
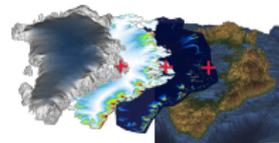


- ▶ 2000–2008 mass changes equally split between changes in surface mass balance and ice discharge (van den Broeke et al, 2009)
- ▶ but simulated ice discharge is nearly constant
- ▶ why do we get such a good agreement with observed mass loss?



We can get “the right result” for the “wrong reason”

# Surface elevation changes 2003–2009



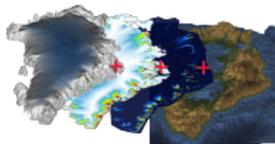
**(a) ICESat**  
(Sørensen et al, 2011)

**(b) ERA init.**

**(c) HadGEM2 init.**

spatially-rich time-series are needed!

# Limitations of hindcasting

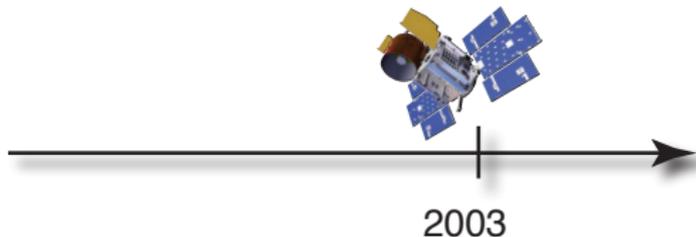


## Theoretical

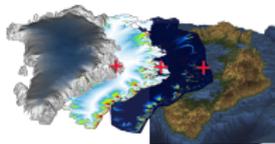
- ▶ The appropriate time-scale for hindcasting is unknown
- ▶ Hindcasts are short (decades) compared to the time-scale associated with changes in energy (thousands of years)
- ▶ Even a hindcast showing good agreement with all available observations may not capture the system's true behavior

## Practical

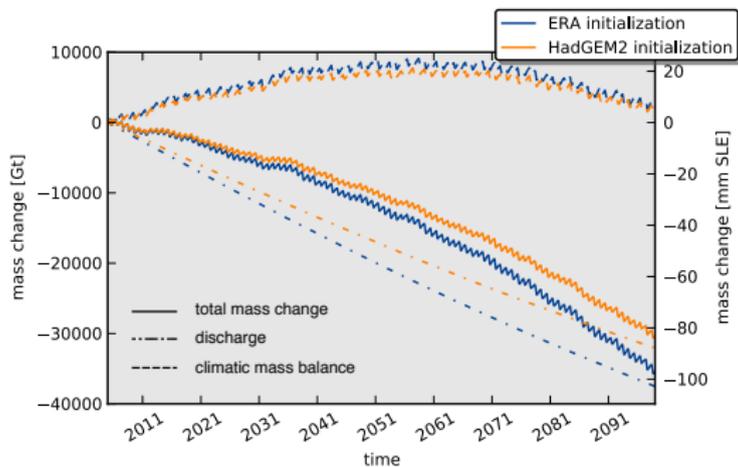
- ▶ Duration of hindcasts is limited by the length of observational records



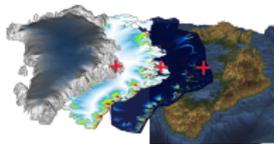
# Forecast



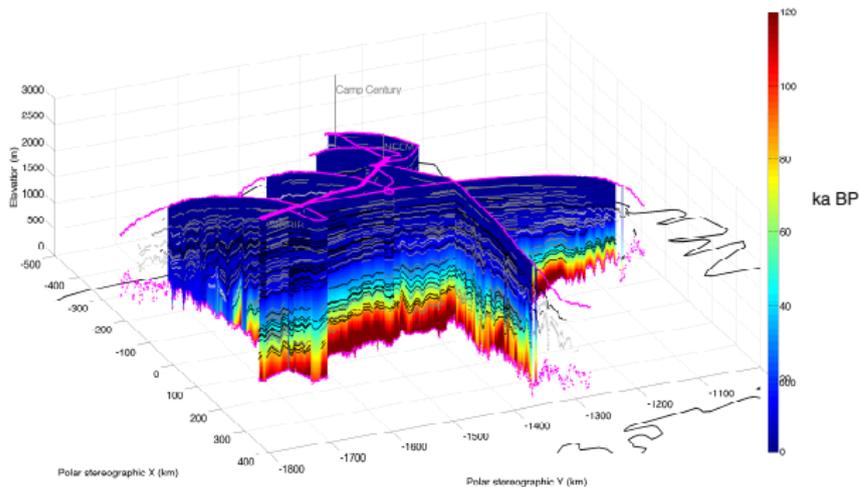
- ▶ only if we're happy with the hindcast



# Outlook: isochrones

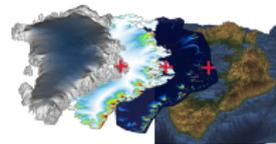


- ▶ distribution of energy within an ice sheet cannot be measured directly
- ▶ age field has similar time-scales

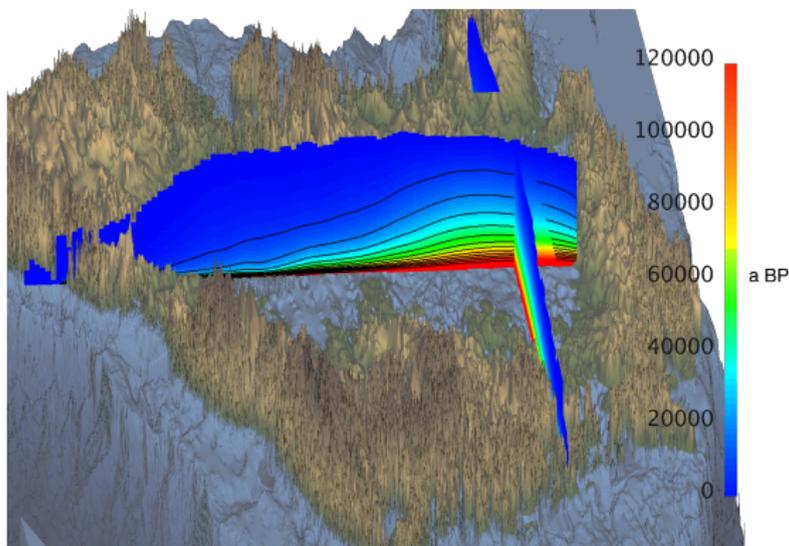


credit: J. MacGregor

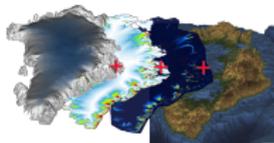
# Outlook: isochrones



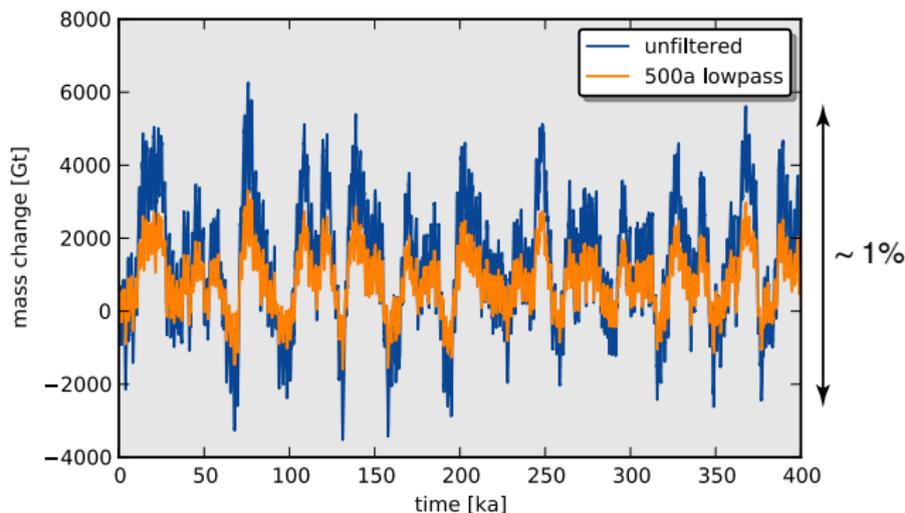
- ▶ distribution of energy within an ice sheet cannot be measured directly
- ▶ age field has similar time-scales



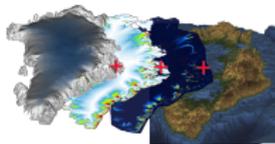
# Outlook: internal variability



- ▶ distribution of energy within an ice sheet cannot be measured directly
- ▶ age field has similar time-scales



# Outlook: statistical frameworks



- ▶ hindcasting may be integrated into comprehensive statistical frameworks to quantify uncertainties in ice sheet evolution due to different sources of model and observation uncertainty

## NASA ROSES Cryosphere

Challenging the Parallel Ice Sheet Model with reproducing the present-day mass loss signal from the Jakobshavn basin, Greenland. PI A. Aschwanden, 2013–2016.